

川东南二叠系龙潭组海-陆过渡相页岩气甜点评价及意义

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摘要:川东南二叠系龙潭组海-陆过渡相沉积是一套潜在的页岩气勘探领域,但目前面临勘探方向不清等问题。综合岩石学、有机地球化学和储层地质学,基于X射线衍射、有机岩石学分析、氩离子抛光扫描电镜、氮气孔隙度测试和联测实验等,分析了川东南地区S1井等重点井龙潭组9个小层中泥页岩有机地化、储层和含气性及其纵向变化特征,优选了页岩气富集的有利层段。研究结果表明,川东南地区龙潭组海-陆过渡相泥页岩沉积厚度大($TOC > 2\%$ 的泥页岩累计厚度50.11 m)、黏土矿物含量高(平均值42.69%),以煤层、炭质页岩、泥页岩、粉砂质泥岩、含灰/灰质泥岩和泥质灰岩为主。泥页岩 TOC 均值为4.63%,孔隙度均值大于2.0%,且都具有煤>炭质页岩>泥页岩>粉砂质泥岩>含灰/灰质泥岩的变化趋势。有机质类型以Ⅲ型为主, R_o 平均为2.07%,储层以黏土矿物孔和微裂缝为主。②—④小层、⑥小层和⑧小层的煤层和泥页岩组合具有“厚度大、有机质丰度高、孔隙度高和含气性好”的特征,是海-陆过渡相页岩气勘探的有利目标层段。

关键词:泥页岩;有机地化;海-陆过渡相;页岩气;龙潭组;二叠系;川东南地区

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Sweet spot assessment and its significance for the marine-continental transitional shale gas of Permian Longtan Fm. in southeastern Sichuan Basin

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Abstract: The marine-continental transitional deposits of Permian Longtan Formation in southeastern Sichuan Basin have been targeted for shale gas but without any significant success. By resorting to analyses on petrology, organic geochemistry, and geology as well as the results of X-ray diffraction, organic petrology analysis, argon ion milling scanning electron microscopy observation, helium porosity testing, and joint measurement, the organic geochemical characteristics, reservoir characteristics, gas-bearing property and their vertical variations of 9 clay mineral-rich (42.69% on average) layers in the Longtan Formation from some key wells (including Well S1) in southeastern Sichuan Basin were studied to locate potential shale gas enrichment intervals. The results show that the marine-continental transitional shale deposits in the Longtan Formation are thick (the cumulative thickness of shale with $TOC > 2\%$ is 50.11 m) and mainly composed of coal seams, carbonaceous shale, shales, silty mudstone, calcareous mudstones and argillaceous limestone. Their average TOC is 4.63% and average porosity is greater than 2.0%, both showing a decreasing trend from coal seam, shale, silty mudstone to calcareous mudstones. The organic matter is mainly of type Ⅲ, with an average R_o of 2.07%. Reservoirs are dominated by clay mineral pores and micro fractures. The combinations of coal with organic-rich shale within layers ② to ④, ⑥ and ⑧ characterized by large thickness, high organic matter

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abundance, high porosity and gas-bearing property are deemed as potential exploration targets for making shale gas discoveries.

Key words: shale, organic geochemistry, marine-continental transitional facies, shale gas, Longtan Formation, Permian, southeastern Sichuan Basin

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继五峰组-龙马溪组页岩气商业开发之后,四川盆地页岩气勘探逐渐向陆相和海-陆过渡相泥页岩转移。四川盆地上二叠统龙潭组海-陆过渡相泥页岩分布范围广、资源潜力大,引起了学者和业界的广泛关注^[1-5]。目前四川盆地龙潭组海-陆过渡相页岩气勘探已取得了积极进展。达县-宣汉地区8口井在龙潭组钻遇良好气测显示,其中明1井在龙潭组4 952~4 970 m深度射孔测试获日产气 $3.02 \times 10^4 \sim 3.85 \times 10^4 \text{ m}^3$,赤水-官渡地区林页1井在龙潭组含气量最高为 $33.78 \text{ m}^3/\text{t}$ 。2018年6月,在丁页2—丁页5井区部署实施中国石化第一口海-陆过渡相页岩气专探井——S1井,共钻遇102.50 m/26层油气显示,其中龙潭组21.00 m/12层,全烃含量最高(23.987%)。

龙潭组海-陆过渡相泥页岩总体沉积于伸展背景,晚二叠世峨眉山玄武岩大规模喷发导致了四川盆地龙潭组出现了沉积分异的格局,横向上河流、滨岸-沼泽、潮坪-潟湖、浅水陆棚、深水陆棚等同期异相,纵向上岩性及其组合复杂多变^[6-8]。其中潮坪-潟湖相沉积主要发育煤层、粉砂质泥岩、泥页岩和灰质泥岩等。前人研究表明四川盆地龙潭组海-陆过渡相泥页岩厚度较大、分布广且稳定,纵向发育多套富有机质泥页岩且与煤层共生。泥页岩有机质丰度高,镜质体反射率(R_o)分布于1.6%~3.0%,已进入大量生气阶段;有机质类型主要为Ⅲ型-Ⅱ₂型干酪根。有机显微组分是制约富有机质泥页岩有机质孔隙发育的重要因素,有机质丰度是影响泥页岩吸附能力和含气性的主要因素^[9-11]。龙潭组海-陆过渡相页岩气勘探和研究处于初级阶段,页岩气富集条件及纵向变化规律尚不明确,直接制约了页岩气勘探层段优选。本次研究以川东南地区S1井为例,综合岩石学、有机地球化学和储层地质学,基于X射线衍射、有机岩石学分析、氦离子抛光扫描电镜、氦气孔隙度测试和联测实验等,分析了龙潭组海-陆过渡相泥页岩有机地化、储层和含气性及其纵向变化特征,优选了页岩气富集的有利层段,以期为四川盆地海-陆过渡相页岩气勘探实践和理论研究奠定基础。

1 地质概况

中二叠世晚期东吴运动导致扬子准地台在经历了中二叠世海盆沉积以后抬升成陆,中二叠统与上二叠统在四川盆地广大地区呈假整合接触,上二叠统底部出现含煤的滨岸-沼泽-潮间沉积^[10-14](图1)。川东南地区S1井钻井揭示二叠系龙潭组主要发育煤、灰黑色炭质页岩、灰黑色页岩、灰黑色泥岩、深灰色细-粉砂质泥岩、深灰色泥质粉砂岩、深灰色灰质泥岩及深灰色泥质灰岩、深灰色灰岩。根据岩性组合将龙潭组划分为9个小层(图2):①小层厚度为4.31 m,以铝土质泥岩为主,底部与茅口组灰岩不整合接触,偶见植物碎屑化石;②—④小层厚度为17.63 m,以灰黑色炭质泥页岩与煤层为主,见3套厚0.50~1.00 m的煤层,煤层附近炭质页岩发育,下部泥页岩颜色较深,向上颜色变浅;⑤小层厚约6.33 m,底部以深灰色泥质灰岩,中、上部见深灰色泥质粉砂岩夹灰黑色泥页岩,为第一套隔夹层,钙质介壳类生物碎屑局部富集;⑥小层厚约25.96 m,为厚度最大的一套富有机质泥页岩,以灰黑色炭质泥页岩与煤层为主,见4套厚0.40~2.50 m的煤层;⑦小层厚约9.61 m,以深灰色灰岩、泥质灰岩和灰质泥岩为主,为第二套隔夹层,局部钙质介壳类生物碎屑富集;⑧小层厚约6.83 m,以灰黑色炭质泥页岩为主,夹1套煤层(单层厚1.19 m),偶见植物碎屑化石;⑨小层厚约17.00 m,以深灰色灰岩为主,夹薄层黑灰色泥岩。灰黑色炭质页岩、灰黑色页岩及灰黑色泥岩主要发育在②—④小层、⑥小层和⑧小层。

2 样品及实验方法

2.1 样品

本文样品主要来自川东南S1井和DYS1井龙潭组(S1井样品181个、DYS1井样品2个)。S1井采用等间

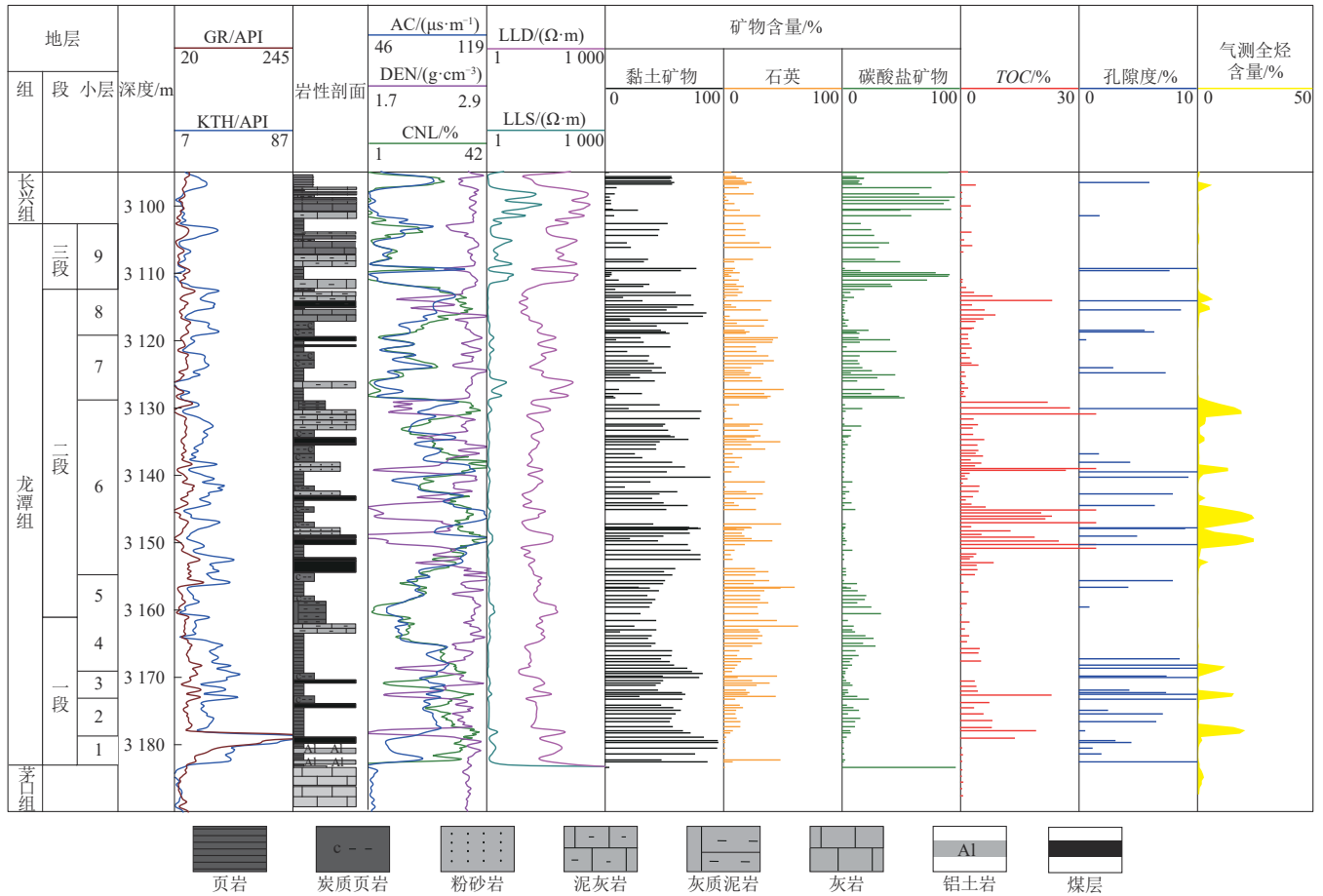


图2 川东南地区S1井龙潭组综合柱状图

Fig. 2 Composite stratigraphic column of the Longtan Formation in Well S1, southeastern Sichuan Basin

征。黏土矿物含量介于1.30%~95.30%，平均值为42.69%，其中②—④小层、⑥小层和⑧小层黏土矿物含量平均值分别平均为54.30%，55.20%和49.50%。石英含量次之，介于0.40%~62.40%，平均值为19.20%。②—④小层、⑥小层和⑧小层石英矿物含量平均值分别为20.50%，20.10%和17.30%。碳酸盐矿物含量较低，平均值为11.07%，②—④小层、⑥小层和⑧小层碳酸盐矿物含量平均值分别为8.60%，3.30%和2.10%(图2)。

4 地球化学特征

4.1 有机质丰度

实测S1井181个样品证实龙潭组泥页岩的有机质丰度变化较大，总有机碳含量(TOC)介于2.06%~38.12%，平均值为4.63%。各小层的TOC差异明显(图2)，除①小层之外，其余各小层TOC均大于1.50%，TOC大于2.00%的泥页岩累计厚度50.11m。

其中⑥小层TOC大于2.00%的泥页岩厚度最大，为17.10m。②—④小层、⑥小层和⑧小层TOC值较高，②—④小层实测22个样品的TOC为2.22%~38.12%，平均值为5.98%；⑥小层实测27个样品的TOC为2.50%~12.88%，平均值为4.47%；⑧小层实测12个样品的TOC为2.21%~16.86%，平均值为5.19%。由此表明，②—④小层、⑥小层和⑧小层泥页岩具有良好的页岩气富集物质基础(图3a)。此外，所有岩性的TOC普遍大于2.00%，但不同岩性的TOC差异较大，煤层TOC平均值为65.93%，炭质泥页岩TOC平均值为16.59%，泥页岩TOC平均值为2.44%，粉砂质泥岩TOC平均值为3.65%，含灰/灰质泥岩TOC平均值为2.11%，即，除了煤层以外，炭质泥页岩和泥页岩的页岩气富集条件较好，粉砂质泥岩和含灰/灰质泥岩次之(图3b)。

4.2 有机显微组分

受沉积环境的影响，川东南地区龙潭组有机母质主要来源于陆源高等植物输入，以Ⅲ—Ⅱ₂型干酪根为

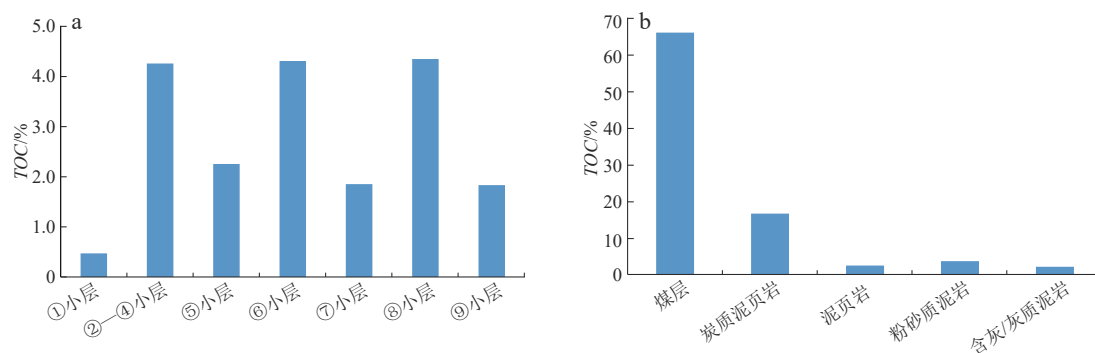


图3 川东南地区S1井龙潭组各小层样品平均TOC(a)及不同岩性样品TOC(b)对比

Fig. 3 Comparison of average TOCs between different layers (a) and different lithologies (b) of the Longtan Formation in Well S1, southeastern Sichuan Basin

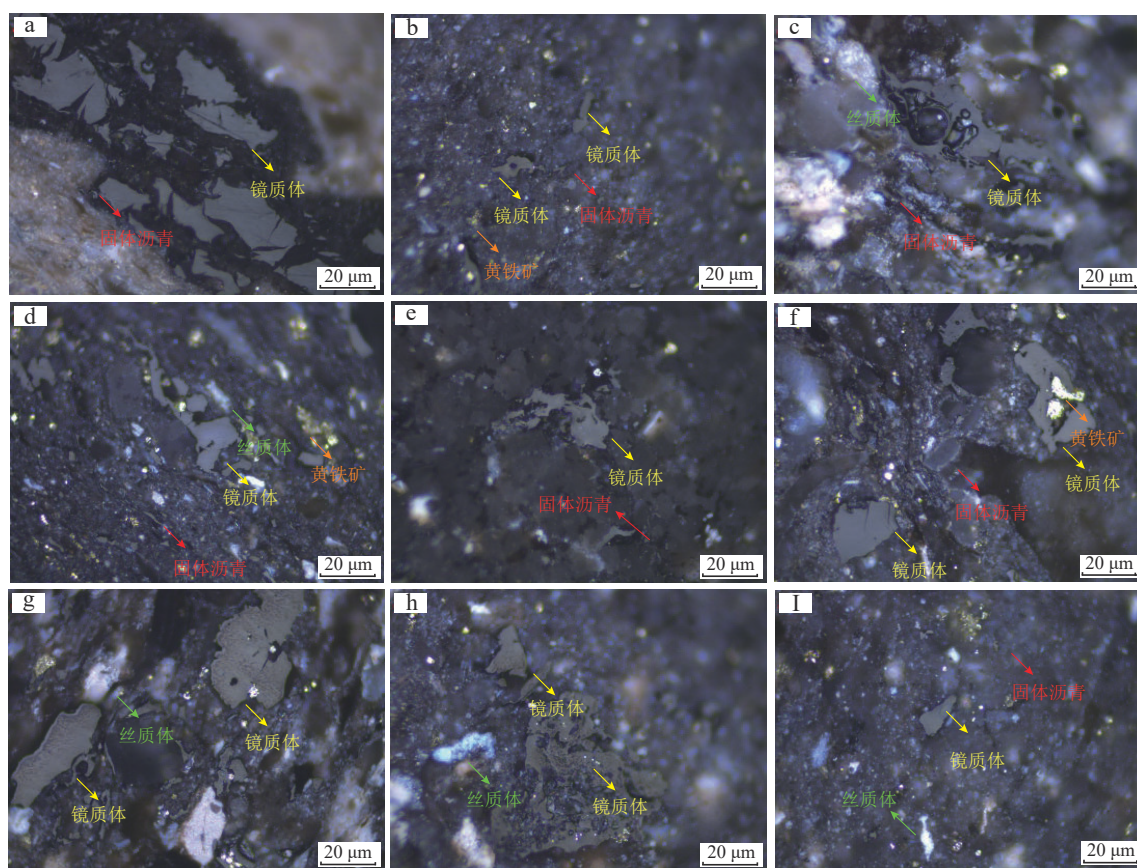


图4 川东南地区S1井龙潭组泥页岩全岩显微组分照片

Fig. 4 Organic macerals of Longtan shale in Well S1, southeastern Sichuan Basin

a. 埋深3 113.40 m, ⑧小层; b. 埋深3 119.69 m, ⑧小层; c. 埋深3 121.73 m, ⑦小层; d. 埋深3 124.90 m, ⑦小层; e. 埋深3 137.50 m, ⑥小层; f. 埋深3 147.35 m, ⑥小层; g. 埋深3 162.52 m, ④小层; h. 埋深3 164.33 m, ④小层; i. 埋深3 174.26 m, ②小层

主^[9]。龙潭组泥页岩有机显微组分以镜质体为主,镜质体含量约为63.93%,呈灰色-灰白色,主要为均质镜质体和无结构镜质体,丝质体次之,反光下呈灰白色-亮白色,且具有微凸起。固体沥青充填裂缝或者孔隙中,形态具有一定的棱角状。龙潭组泥页岩以Ⅲ型干酪根为主,实测干酪根碳同位素值 $\delta^{13}\text{C}$ 介于-24.0‰~

-22.1‰。综合判断认为,龙潭组以腐殖型有机质为主,有机母质主要为陆源高等植物。②—④小层顶部见少量固体沥青,⑥小层中上部见固体沥青,固体沥青含量比②—④小层略有增多,⑧小层固体沥青进一步增多,由此反映了该地区龙潭组泥页岩具有较好的生气能力(图4),为一套优质气源岩。

表1 川东南地区S1井龙潭组实测样品镜质体反射率数据
Table 1 Measured vitrinite reflectance of samples from the Longtan Formation in Well S1, southeastern Sichuan Basin

深度/m	岩性	$R_o/\%$	测点数/个
3 095.98	黑色页岩	1.86	32
3 109.72	黑灰色页岩	1.93	23
3 114.20	黑色炭质页岩	1.99	40
3 130.21	煤层	2.08	29
3 143.55	灰黑色页岩	2.09	32
3 166.90	黑灰色泥岩	2.13	20
3 169.95	灰黑色炭质页岩	2.16	30
3 178.38	煤层	2.20	31
3 179.52	黑色炭质页岩	2.21	32

4.3 热演化程度

较高的热演化程度是热成因页岩气生成和富集的必要条件,一方面高热演化程度可以为页岩气富集提供充足物源,另一方面高热演化程度也是有机质孔发育的前提条件^[15]。为了表征龙潭组烃源岩热演化程度,本次研究实测了9个样品的镜质体反射率(R_o),每个样品的测点数均大于或等于20(表1)。统计结果表明,龙潭组样品 R_o 主体分布于1.86%~2.21%,平均值为2.07%,现今热演化程度较高,处于主生气期,对页岩气的富集十分有利。

5 储层及含气性特征

5.1 储层物性特征

实测样品孔隙度结果表明,龙潭组各小层的储层孔隙度均大于2.00%,具有良好的储层物性条件。与

TOC相似,各小层孔隙度变化较大(图2),②—④小层、⑥小层和⑧小层储层物性较高,孔隙度均值分别为7.37%、8.41%和9.74%(图5a)。此外,不同岩性的孔隙度变化也具有如下趋势:煤层>炭质泥页岩>泥页岩>粉砂质泥岩>含灰/灰质泥岩,其平均孔隙度依次为10.91%、8.63%、7.00%、4.26%和4.01%,即除了煤层以外,炭质泥页岩和泥页岩的储层物性最好(图5b)。

此外,龙潭组样品覆压孔隙度数据表明,虽然不同岩性样品的初始孔隙度存在差异,但是覆压条件下所有岩性样品的孔隙度没有明显降低,当压力为50 MPa时样品孔隙度仅减少了0.59%~1.45%,与五峰组-龙马溪组硅质页岩基本相当,说明储层具有一定的抗压实能力(图6)。

5.2 孔隙类型及孔径特征

扫描电镜观察结果表明S1井龙潭组黏土矿物层间或晶间发育大量孔隙(图7),孔隙主要为片状、线状或狭缝状,局部狭缝状孔隙可形成微裂缝。孔隙长度和宽度变化较大,孔径主要介于几十到几百纳米,缝宽可达457.9 nm,孔隙多被有机质充填或半充填。孔隙主要发育在伊/蒙混层间,部分微孔隙发育在伊利石、高岭石和绿泥石中。

龙潭组高成熟-过成熟泥页岩处于晚成岩阶段,黏土矿物主要为伊利石,局部可见绿泥石或伊/蒙混层^[16]。缠绕或扭曲状伊利石片层之间发育丰富的楔形或狭缝状孔缝,孔缝的宽度变化较大,延伸程度不一,局部可形成复杂缝网^[17]。绿泥石片层平行排列紧密,局部板片间见狭缝,缝宽变化较大。局部发育的伊/蒙混层中可见席状或板状孔隙,孔隙延伸性较差(图7)。

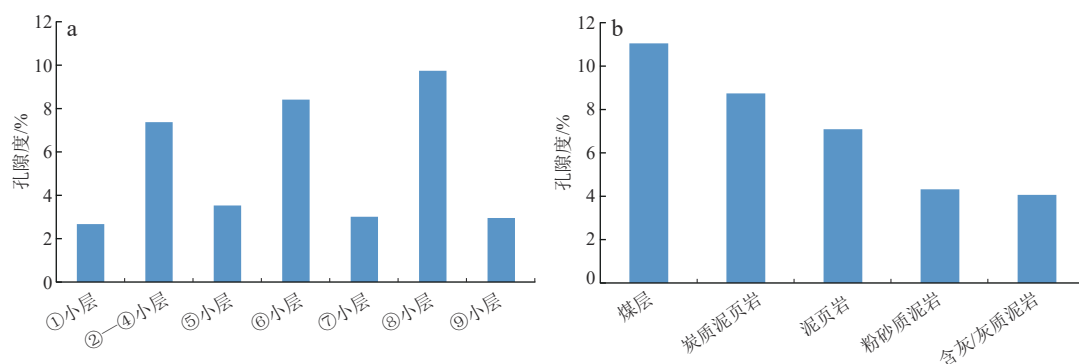


图5 川东南地区S1井龙潭组各小层样品平均孔隙度(a)及不同岩性样品孔隙度(b)对比

Fig. 5 Comparison of average porosity between different layers (a) and different lithologies (b) of the Longtan Formation in Well S1 in the southeast of the Sichuan Basin

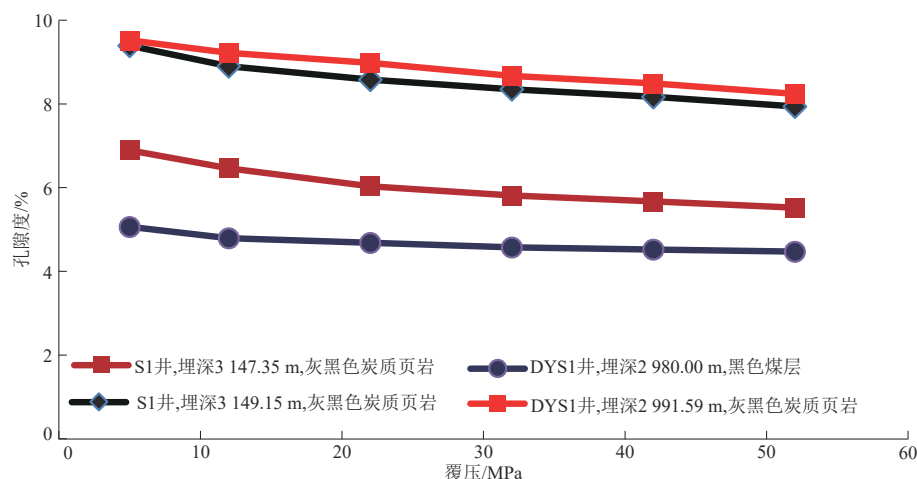


图6 川东南地区龙潭组样品孔隙度随覆压变化趋势

Fig. 6 Variation of porosity with overburden pressure of the samples from the Longtan Formation, southeastern Sichuan Basin

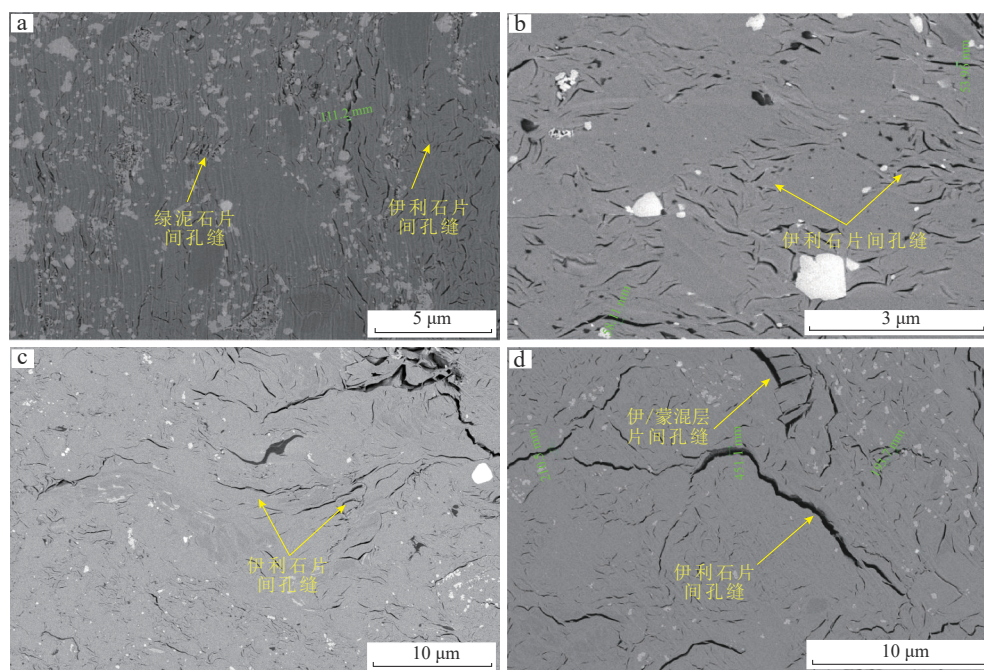


图7 川东南地区S1井龙潭组泥页岩黏土矿物孔隙SEM照片

Fig. 7 SEM images showing clay mineral pores in shale/mudstone of Longtan Formation in Well S1, southeastern Sichuan Basin

a. 埋深3 148.02 m, 灰黑色页岩; b, c. 埋深3 154.44 m, 灰黑色泥岩; d. 埋深3 170.61 m, 灰黑色页岩

龙潭组泥页岩样品内可见大量的暗色、具有一定颗粒形态(多为条带状或块状)的有机质。受有机显微组分控制,多数有机质边缘多为平滑、表面光洁、质地均匀、结构紧密,内部不发育孔隙^[18-20]。局部发育形状不规则、蜂窝状有机质孔隙,有机质孔孔径变化较大,最大孔径可达百纳米级。部分暗色有机质内发育孤立圆形孔,孔径最大可达900 nm,孔隙内半充填。该类孔隙为原始孔隙或者继承性孔隙,主要发育在高等植物的茎或叶片之内,主要通过继承细胞原始结构而形成,孔隙发育不受热演化作用以及成岩作用的控

制。此外,孔隙无明显的长轴方向,由此推测机械压实作用对孔隙保存影响较小(图8)。

综合对比发现,川东南地区龙潭组泥页岩储层整体以无机矿物孔为主,主要为黏土矿物相关孔隙,与川东地区侏罗系陆相页岩具有一定的相似性^[21-22]。由于上部泥页岩中固体沥青含量有所增加,上部各小层有机质孔隙数量增多。

龙潭组炭质泥页岩和泥页岩储层主要以介孔为主(占比50%~70%),微孔次之(占比20%)。这与扫描电镜下观察到的黏土矿物孔为主的孔隙发育特征

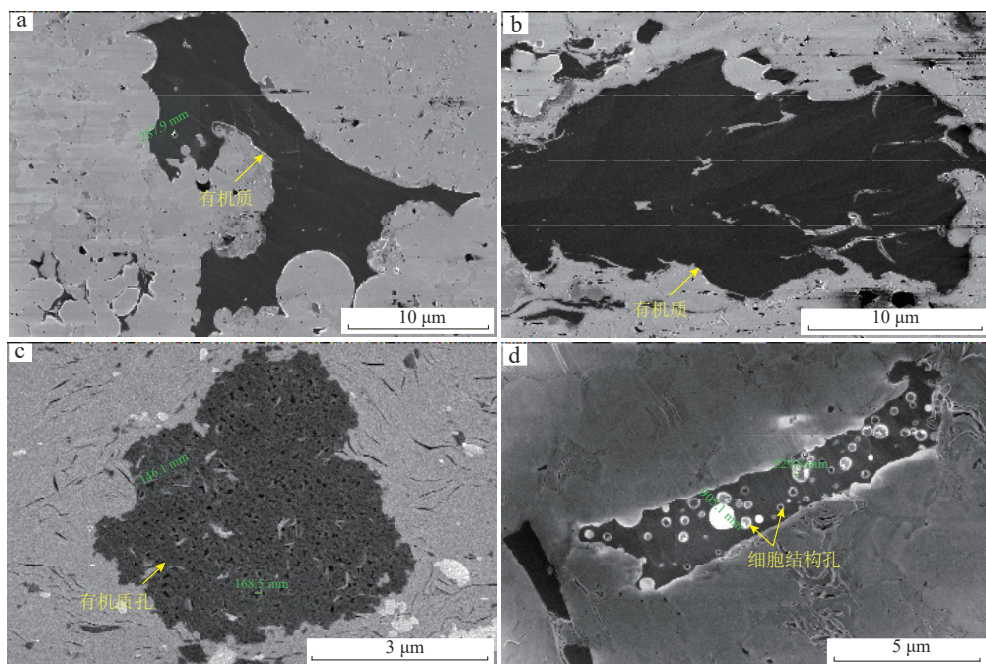


图8 川东南地区S1井龙潭组页岩有机质孔隙SEM照片

Fig. 8 SEM images showing organic pores of shale/mudstone in the Longtan Formation in Well S1, southeastern Sichuan Basin

a. 埋深3 156.75 m, 灰黑色页岩; b. 埋深3 158.56 m, 黑灰色粉砂质页岩; c. 埋深3 155.66 m, 黑灰色泥岩; d. 埋深3 162.52 m, 黑灰色含灰泥岩

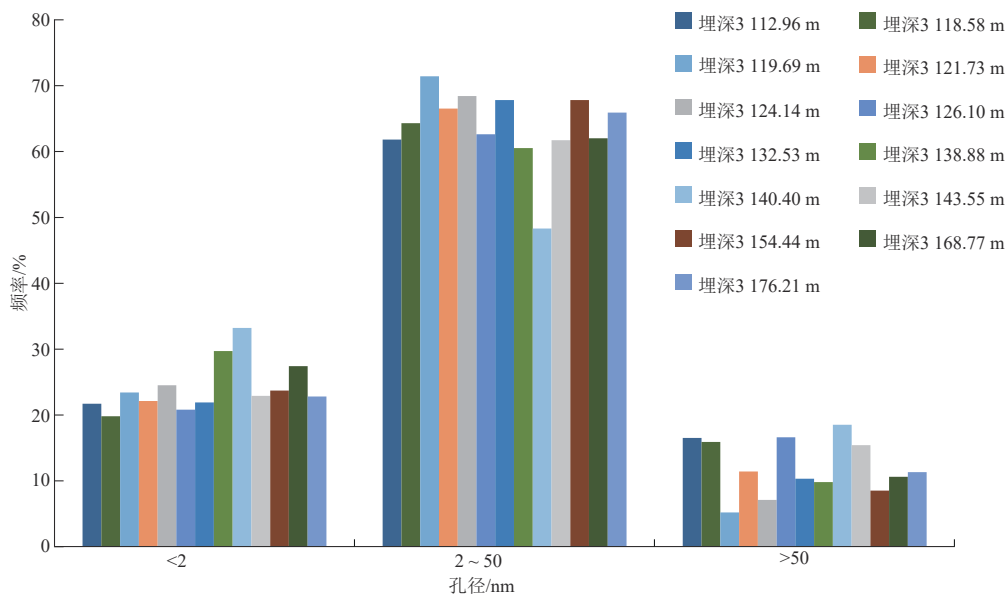


图9 川东南地区S1井龙潭组泥页岩孔径分布频率直方图

Fig. 9 Pore size histogram of shale/mudstone in the Longtan Formation in Well S1, southeastern Sichuan Basin

相吻合。此外,各小层样品的孔径分布无明显的差异(图9)。

5.3 含气性特征

实测含气量数据表明龙潭组潮坪-潟湖相沉积普遍含气,但纵向含气性变化较大,其中,②—④小层、⑥小层和⑧小层的含气性相对较好(图10),含气量分

别为1.81,2.73和2.08 m³/t。该变化趋势与龙潭组TOC和孔隙度的趋势基本一致,说明有机质和储层储集能力对海-陆过渡相泥页岩含气性具有重要的控制作用^[23-24]。虽然龙潭组泥页岩有机质孔发育程度较低,但是高有机质丰度保障了泥页岩的生气潜力。由于黏土矿物存在电荷,水分子与黏土可以紧密结合,导致黏土矿物颗粒表面存在“水膜”,表现为亲水特

表2 川东-川东南地区不同相带页岩气富集条件对比

Table 2 Comparison of shale gas enrichment conditions in various facies zones in eastern and southeastern Sichuan Basin

层系	井名	沉积亚相	TOC/%	干酪根类型	R_o /%	孔隙度/%	孔隙类型	脆性矿物含量/%	黏土矿物含量/%
二叠系龙潭组	S1	潮坪-潟湖	3.84	Ⅲ-Ⅱ ₂	2.07	8.41	黏土矿物层间孔为主,可见有机质孔	44.80	55.20
五峰组-龙马溪组	JY1	深水陆棚	3.52	Ⅱ ₁	2.00~2.50	4.81	有机质孔为主,可见脆性矿物孔	62.40	34.60
侏罗系自流井组东岳庙段	FY10	半深湖-深湖	1.37	Ⅲ-Ⅱ ₂	1.56	4.15	黏土矿物层间孔为主,可见有机质孔	43.82	56.18

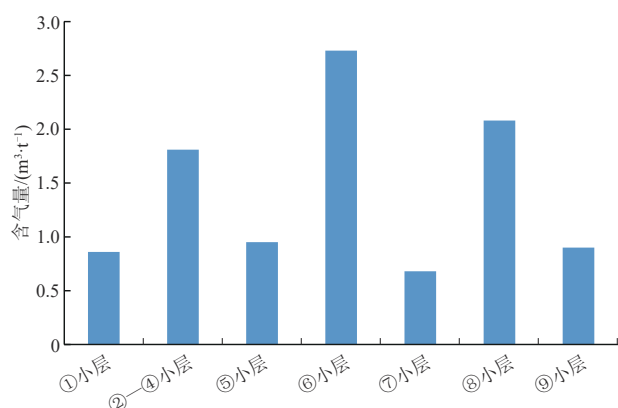


图10 川东南地区S1井龙潭组含气量柱状图

Fig. 10 Bar chart showing gas content of shale/mudstone in the Longtan Formation in Well S1, southeastern Sichuan Basin

征^[24-26]。因此,在地层水存在的情况下,黏土矿物吸附甲烷的表面积降低,影响了吸附能力,从而不利于甲烷富集^[27-34]。钻井揭示龙潭组具有良好的气测显示,也说明龙潭组具有普遍含气的特征,②—④小层、⑥小层和⑧小层的全烃最高值分别为15.373%、23.987%和6.469%,其中煤及与之相邻的高TOC泥页岩含气性最好。

6 地质意义

综上所述可知,研究区龙潭组潮坪-潟湖相沉积具有良好的生烃和成储条件,此外,②—④小层TOC大于2%的泥页岩厚度13.64 m,最大连续厚度5.36 m;⑥小层TOC大于2%的泥页岩厚度17.10 m,最大连续厚度5.12 m;⑧小层TOC大于2%的泥页岩厚度4.76 m,最大连续厚度3.68 m。综合泥页岩厚度、TOC、孔隙度和含气性判断,这3套泥页岩中页岩气资源潜力和勘探前景较好,其中⑥小层页岩气富集条件优越。与川东地区五峰组-龙马溪组海相页岩相比,二叠系龙潭组海-陆过渡相泥页岩具有较高的有机质丰度和

孔隙度,但干酪根类型较差。与川东地区侏罗系自流井组页岩相比,研究区龙潭组潮坪-潟湖相泥页岩具有高TOC、高 R_o 和高孔隙度,说明具有良好的页岩气富集条件(表2)。

7 结论

1) 龙潭组海-陆过渡相沉积以煤层、炭质泥岩、泥页岩、粉砂质泥岩、泥质粉砂岩、含灰/灰质泥岩、泥质灰岩和灰岩为主,TOC大于2%的泥页岩厚度50.11 m,⑥小层TOC大于2%的泥页岩厚度最大,为17.10 m。有机质丰度为2.06%~38.12%,平均为4.63%,具有煤>炭质泥页岩>泥页岩>粉砂质泥岩>灰质泥岩的变化趋势;热演化程度较高, R_o 为1.86%~2.21%,平均为2.07%;有机显微组分主要为镜质体、丝质体、固体沥青,见少量惰屑体、粗粒体,⑥小层和⑧小层较②—④小层固体沥青含量略有增多。

2) 龙潭组海-陆过渡相泥页岩储层孔隙度较高,均值大于2.0%,同样具有煤>炭质泥页岩>泥页岩>粉砂质泥岩>灰质泥岩的变化趋势,储层以黏土矿物孔和微裂缝为主,主要发育伊利石片间孔缝,有机质孔呈局部发育特征,孔隙以介孔为主。

3) 龙潭组海-陆过渡相泥页岩储层②—④小层、⑥小层和⑧小层煤层与泥页岩组合具有“厚度大、有机质丰度高、孔隙度高和含气性好”的特征,是海-陆过渡相页岩气勘探的有利层段。

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